

Evaluation of the Arctic Surface Radiation Budget in CMIP5 models

GOAL →

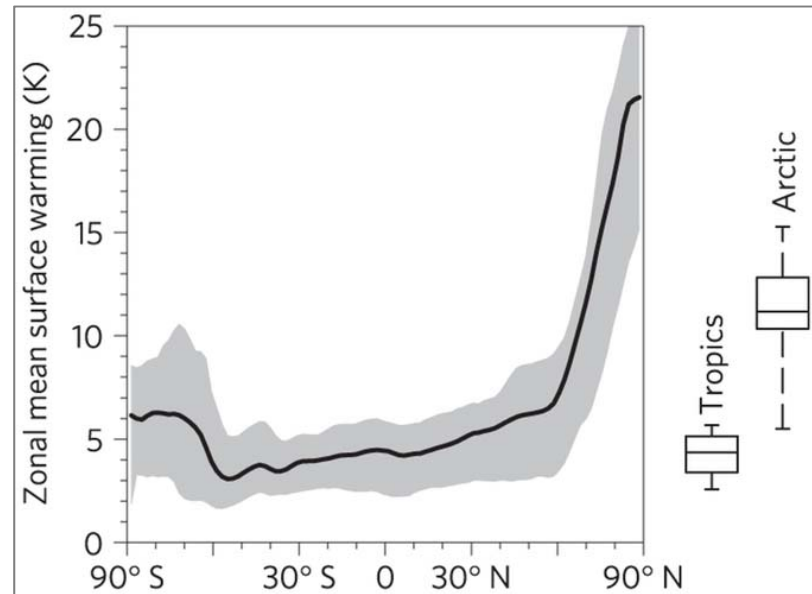
Determine biases in the representation of the Arctic surface radiation budget annual cycle and discover the physical processes that explain the significant spread in projected Arctic warming.

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image credit: www.awi.de

The Arctic climate is rapidly changing

Arctic surface temperature is increasing at a rate outpacing the rest of the globe, and the projected Arctic temperature response to increasing CO₂ is larger than that for the tropics.



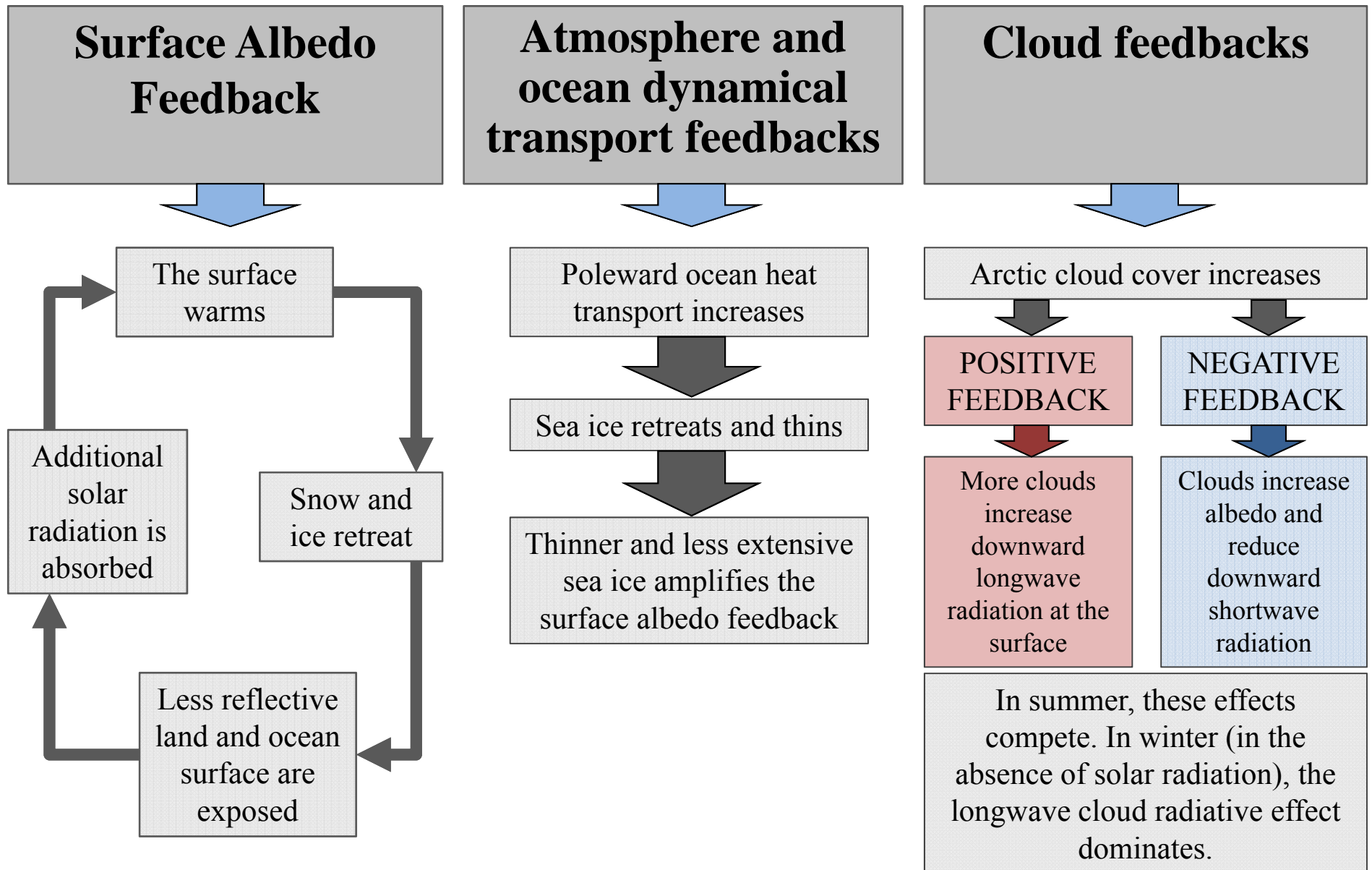
credit: Pithan and Mauritsen 2013

Studying the Arctic climate presents unique challenges.

- The largest intermodel spread in projected surface temperature warming is found in the Arctic.
- Satellite observations are difficult, lack of in-situ measurements

Understanding and reducing intermodel spread in the simulation of the surface energy budget can improve future projections.

Radiative and non-radiative feedback processes lead to polar warming amplification



Use the concept of cloud radiative forcing to evaluate the influence of clouds on shortwave and longwave fluxes at the surface.

$$\text{CRE} = (\text{SW}\downarrow - \text{SW}\downarrow_{\text{clr-sky}}) \cdot (1 - \alpha) + (\text{LW}\downarrow - \text{LW}\downarrow_{\text{clr-sky}})$$



“Cloud Radiative Effect”

where

- $\text{SW}\downarrow, \text{LW}\downarrow$ are all-sky fluxes
- $\text{SW}\downarrow_{\text{clr-sky}}, \text{LW}\downarrow_{\text{clr-sky}}$ are clear-sky fluxes
- α is the albedo calculated using clr-sky sw fluxes, $\text{SW}\uparrow_{\text{clr-sky}}/\text{SW}\downarrow_{\text{clr-sky}}$

Terms in the equation represent cloud influence on solar and infrared radiation

$$(\text{SW}\downarrow - \text{SW}\downarrow_{\text{clr-sky}}) \cdot (1 - \alpha)$$



Shortwave cloud radiative forcing (SW CRE)

Usually negative because downwelling solar flux decreases with the presence of clouds

Magnitude of SW CRE is smaller over a white surface than over ocean

$$(\text{LW}\downarrow - \text{LW}\downarrow_{\text{clr-sky}})$$

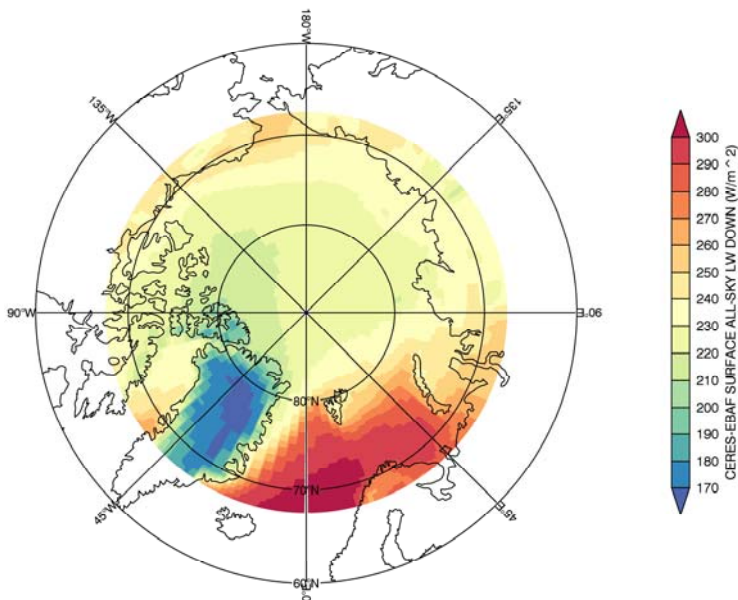
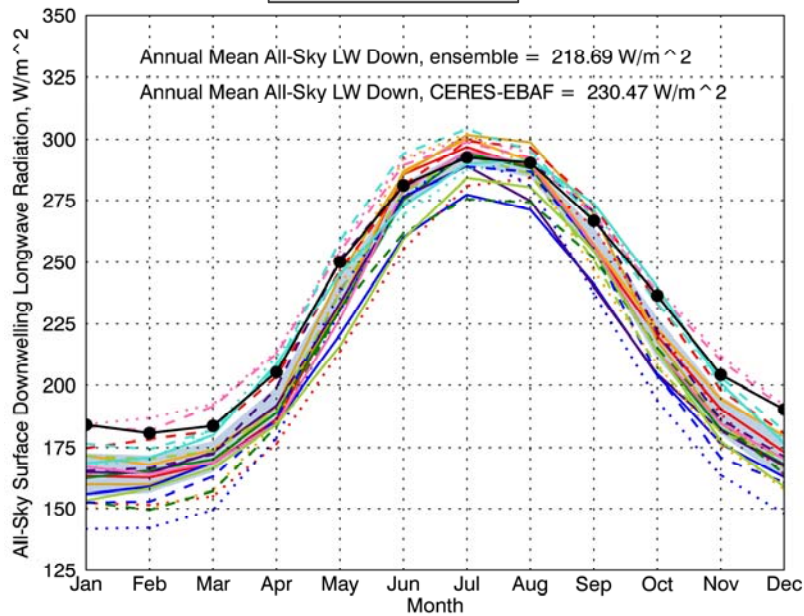


Longwave cloud radiative forcing (LW CRE)

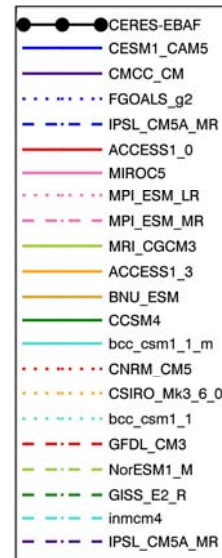
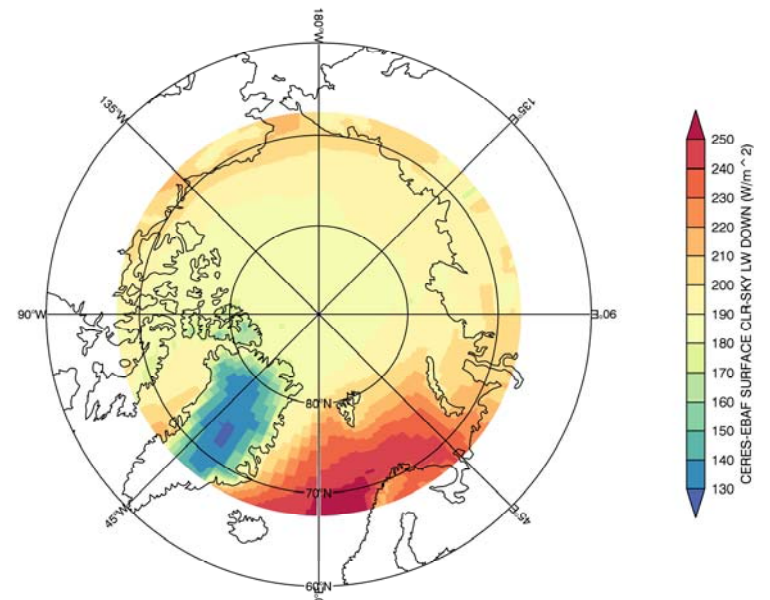
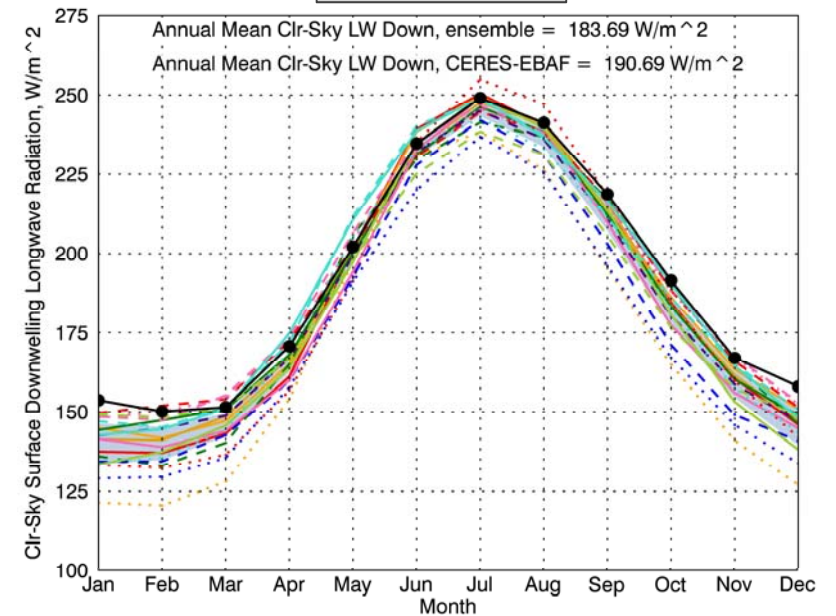
Usually positive because downwelling longwave radiation increases with the presence of clouds

Longwave Surface Fluxes

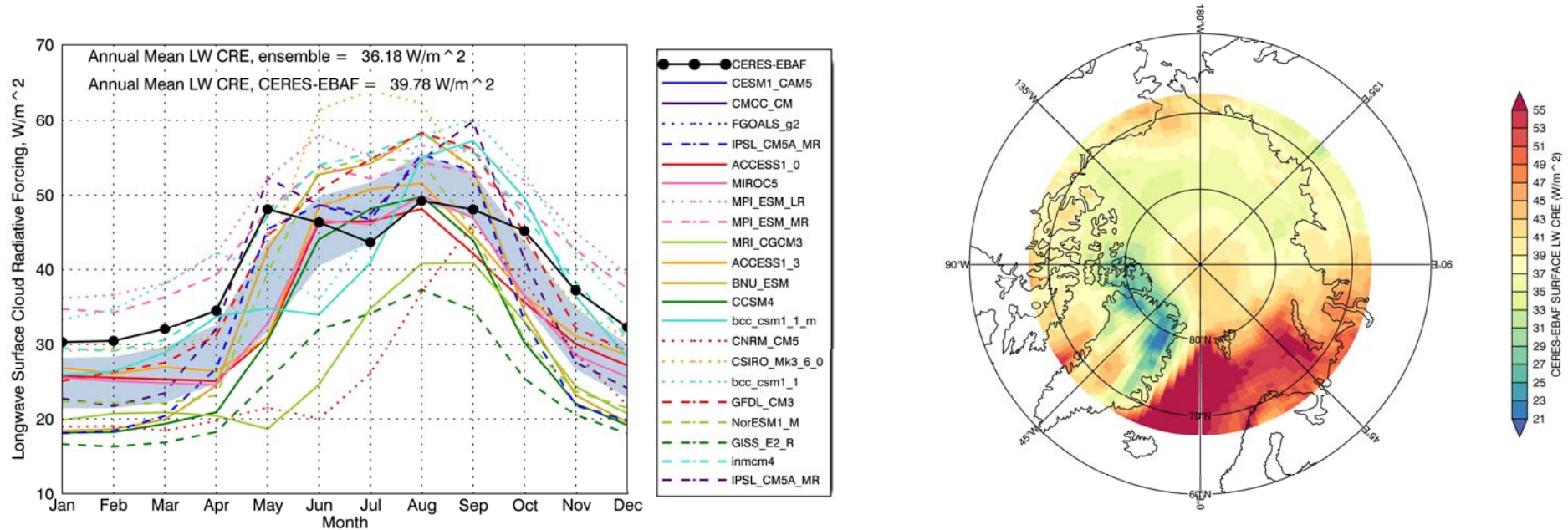
All-Sky



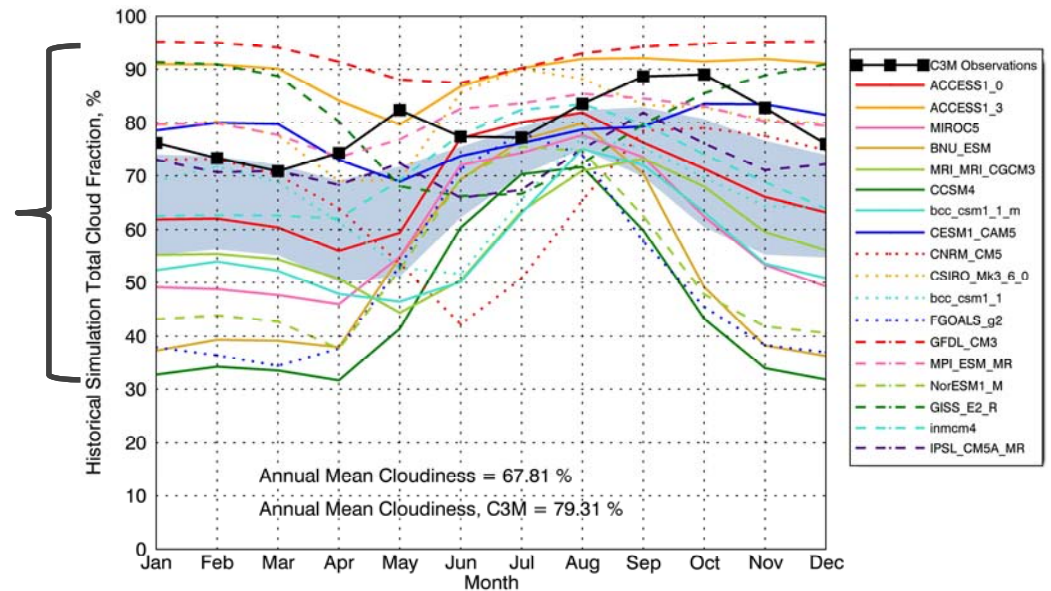
Clr-Sky



Longwave Cloud Radiative Effect



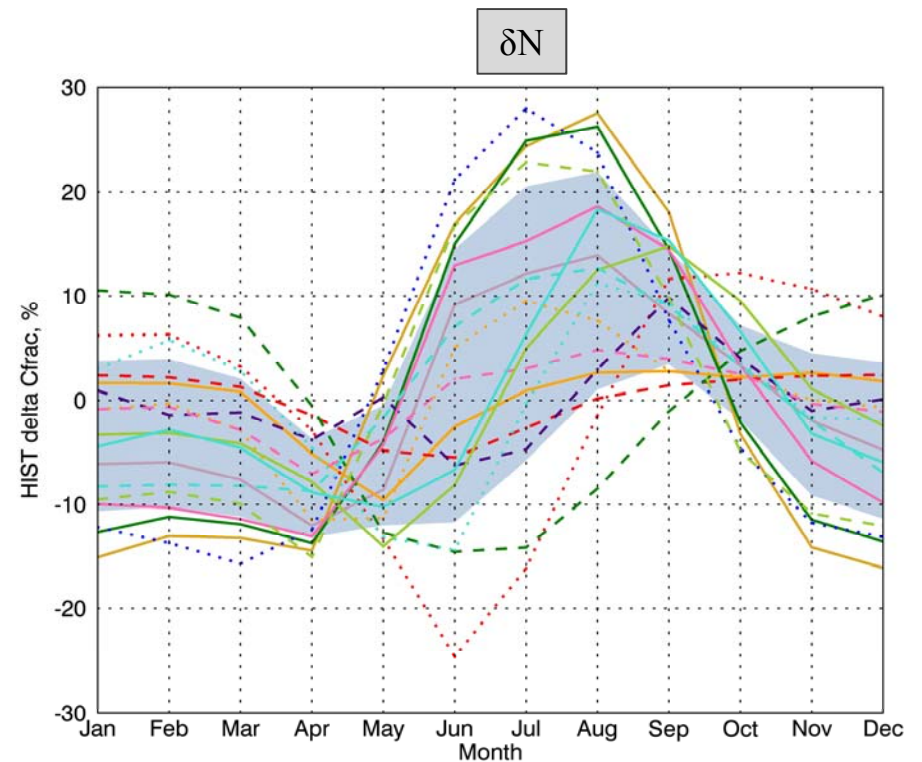
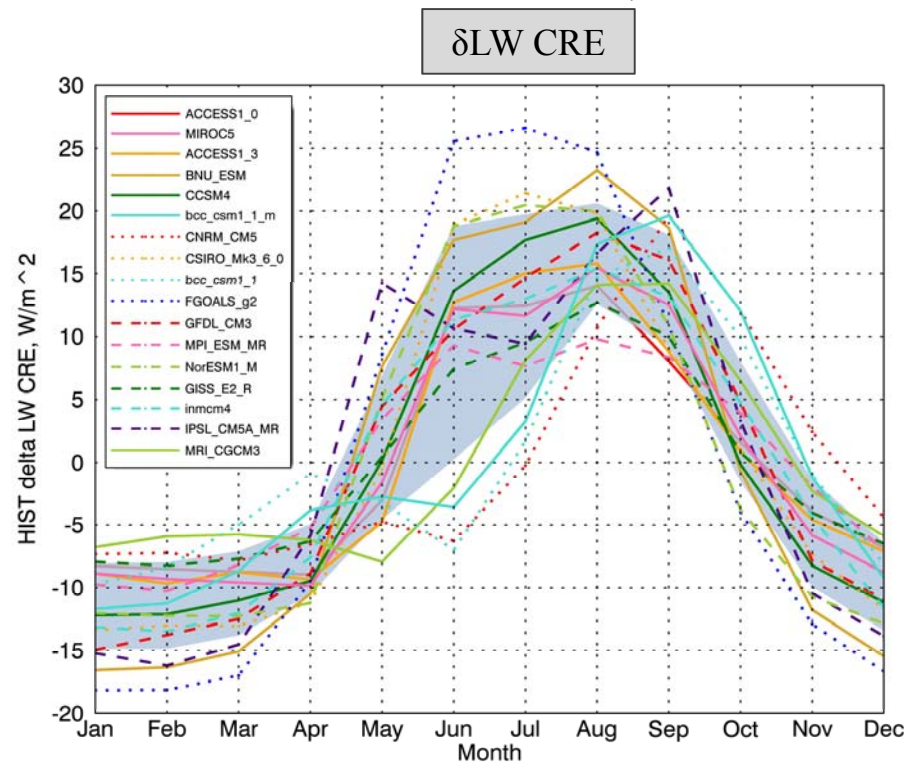
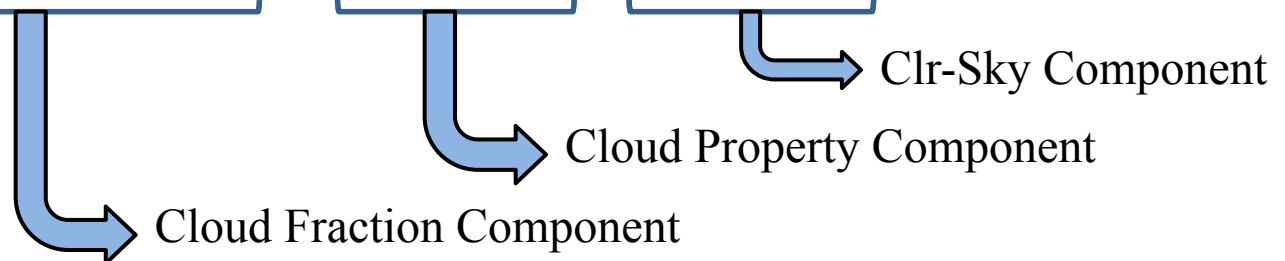
The large discrepancy in wintertime cloudiness is due to the representation of low clouds (Karlsson 2011)



What causes differences in LW CRE?

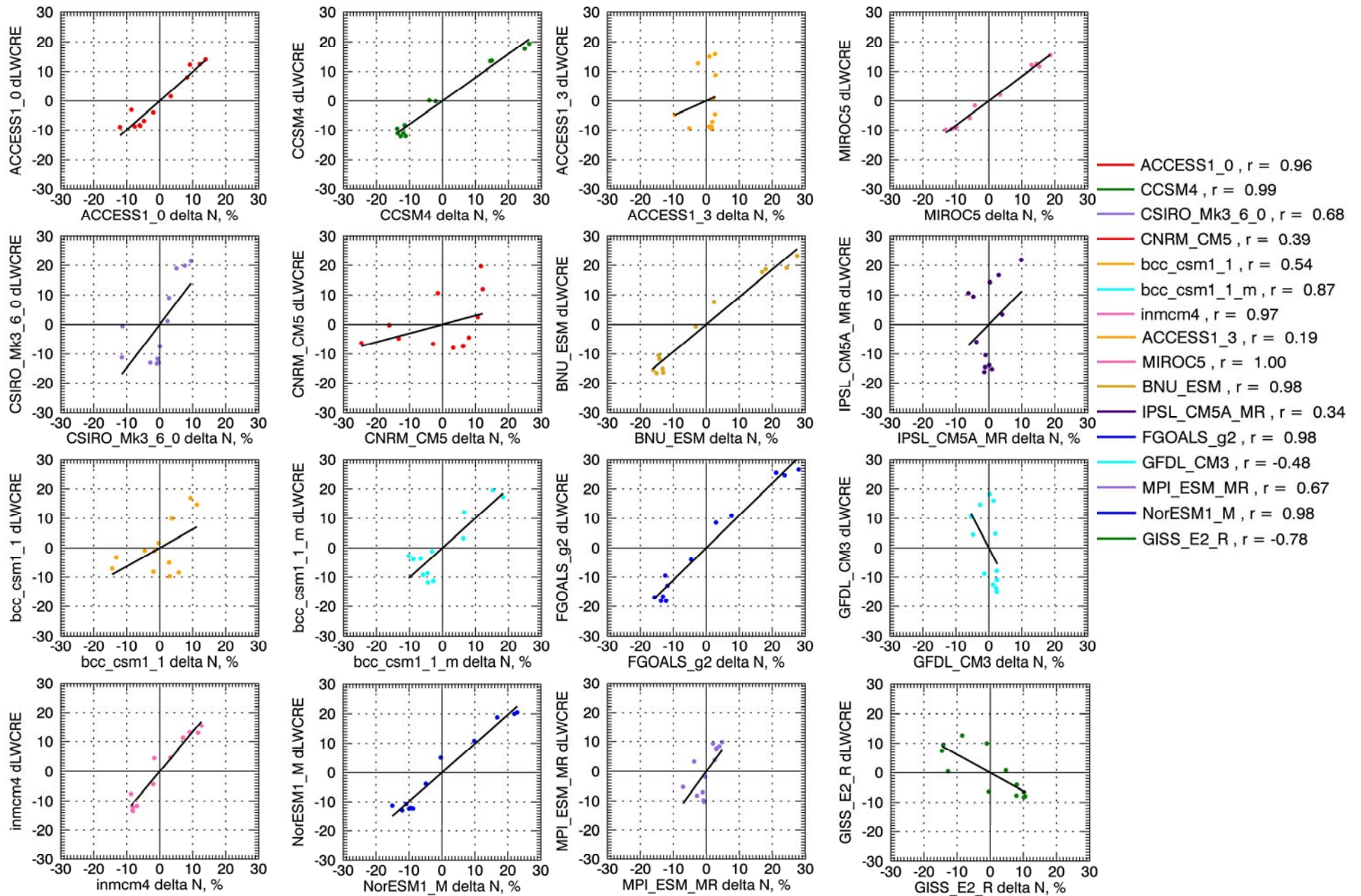
$$\text{LW CRE} = \text{LW} \downarrow_{\text{all}} - \text{LW} \downarrow_{\text{clr}} = N(F \downarrow_{\text{cld}, \text{lw}} - F \downarrow_{\text{clr}, \text{lw}})$$

$$\delta \text{LW CRE} = \underbrace{\delta N(F \downarrow_{\text{cld}, \text{lw}} - F \downarrow_{\text{clr}, \text{lw}})}_{\text{Cloud Fraction Component}} + \underbrace{N \times \delta F \downarrow_{\text{cld}, \text{lw}}}_{\text{Cloud Property Component}} - \underbrace{N \times \delta F \downarrow_{\text{clr}, \text{lw}}}_{\text{Clr-Sky Component}}$$



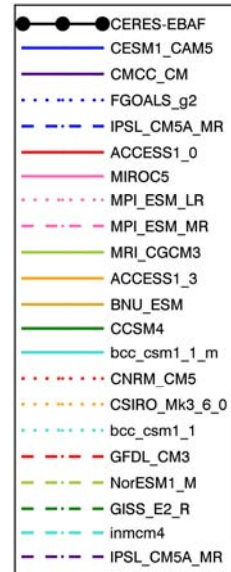
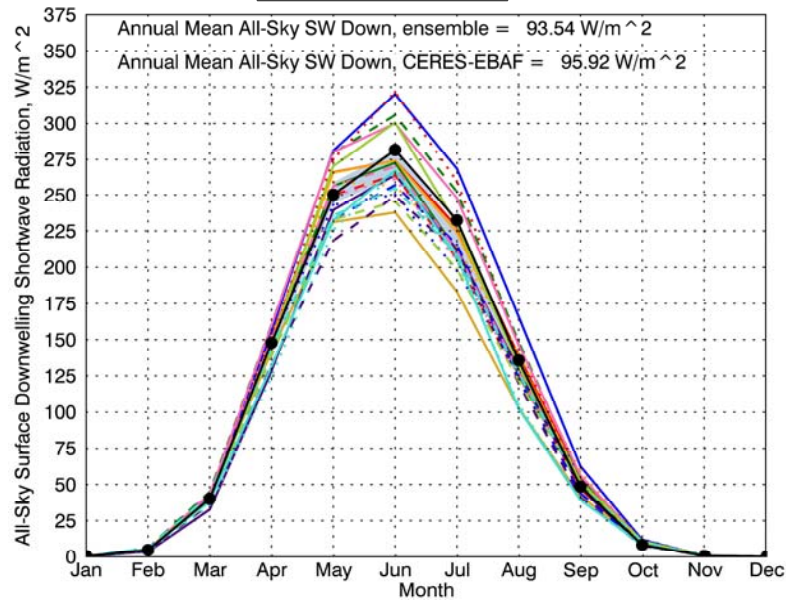
(grey shaded region is the ensemble mean +/- one standard deviation)

For some models, changes in LW CRE are closely coupled to changes in cloud fraction

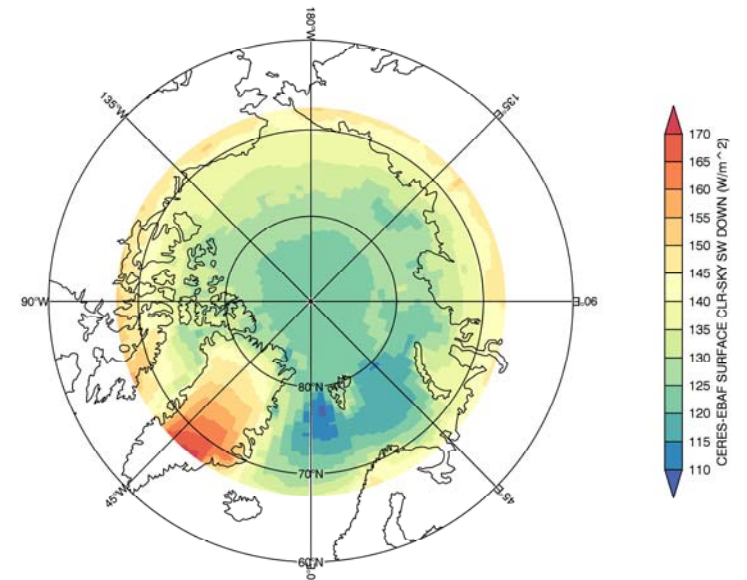
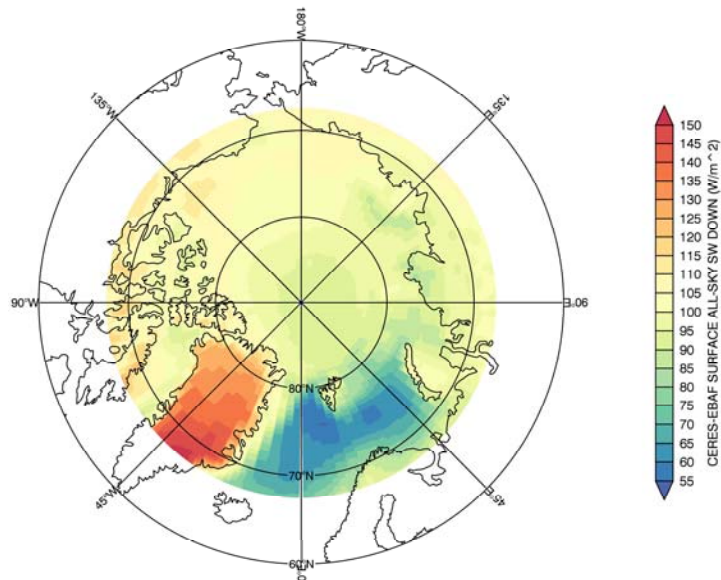
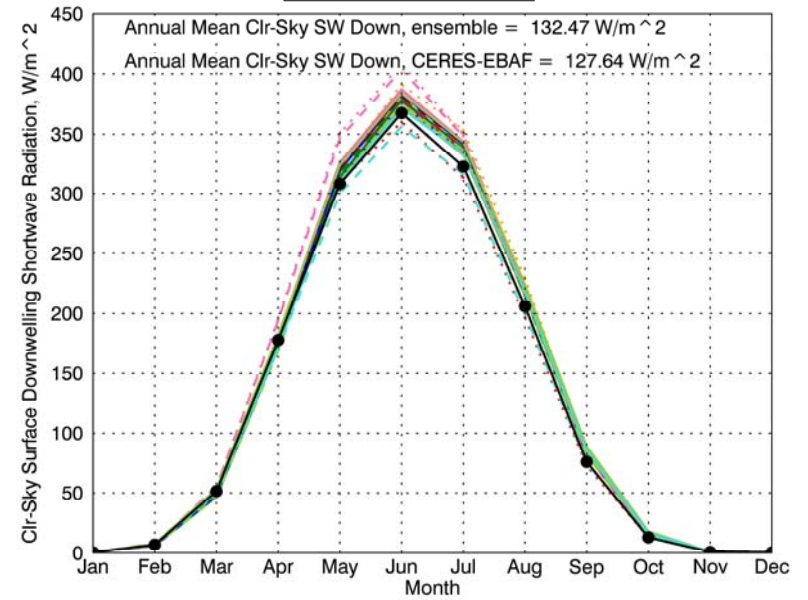


Shortwave Surface Fluxes

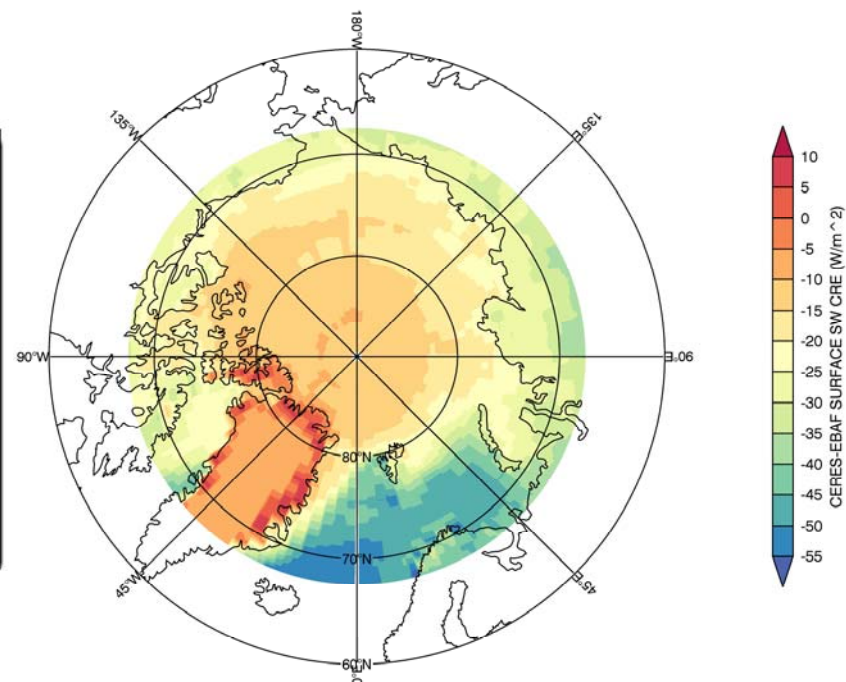
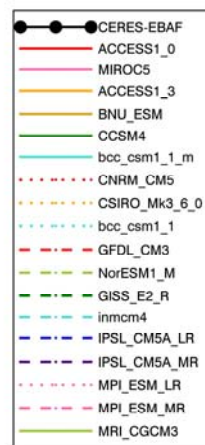
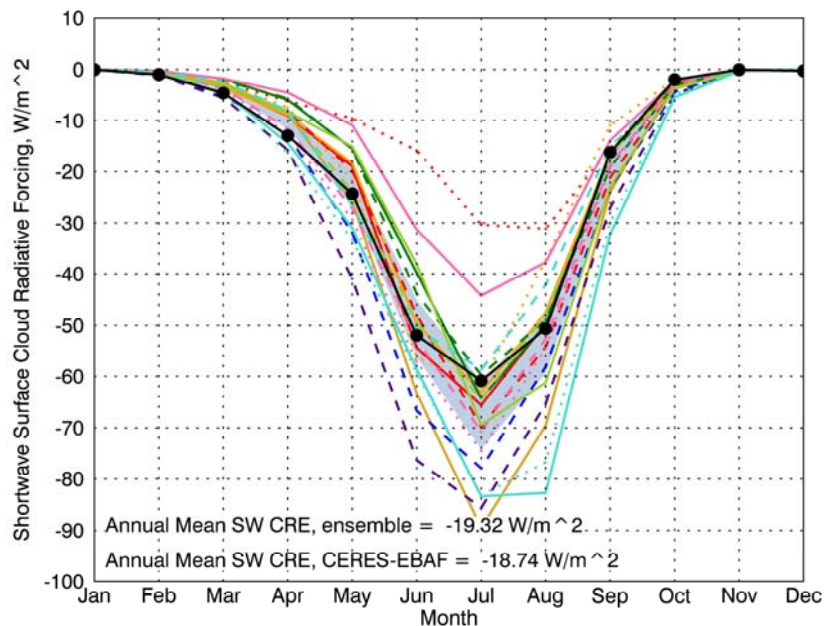
All-Sky



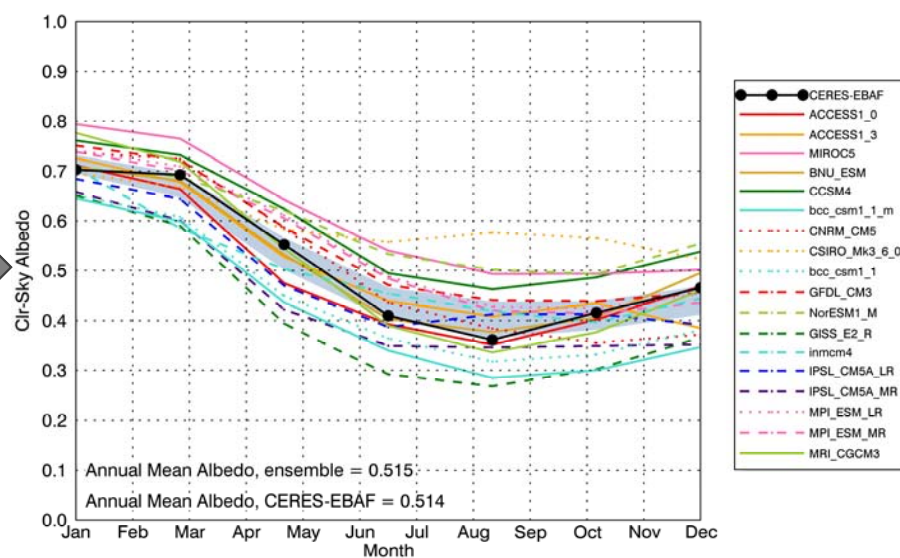
Clr-Sky



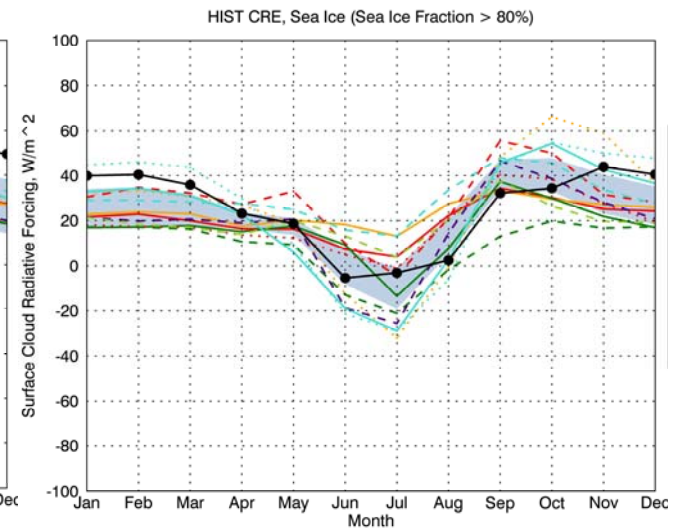
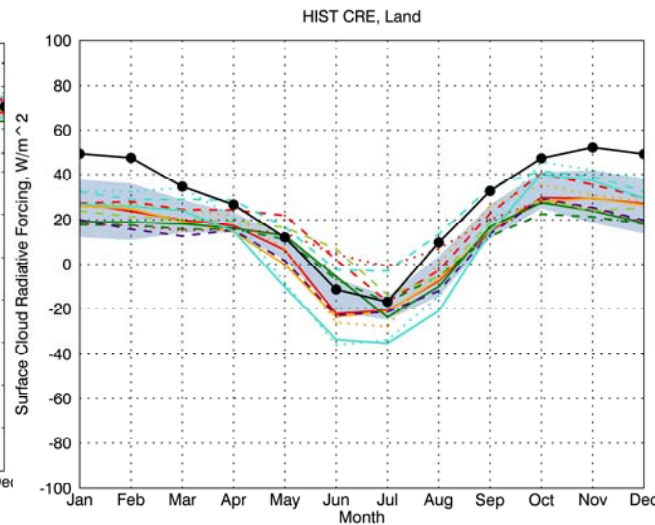
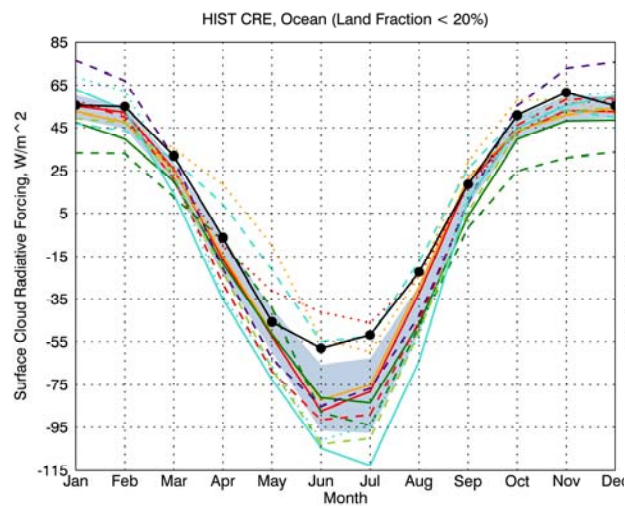
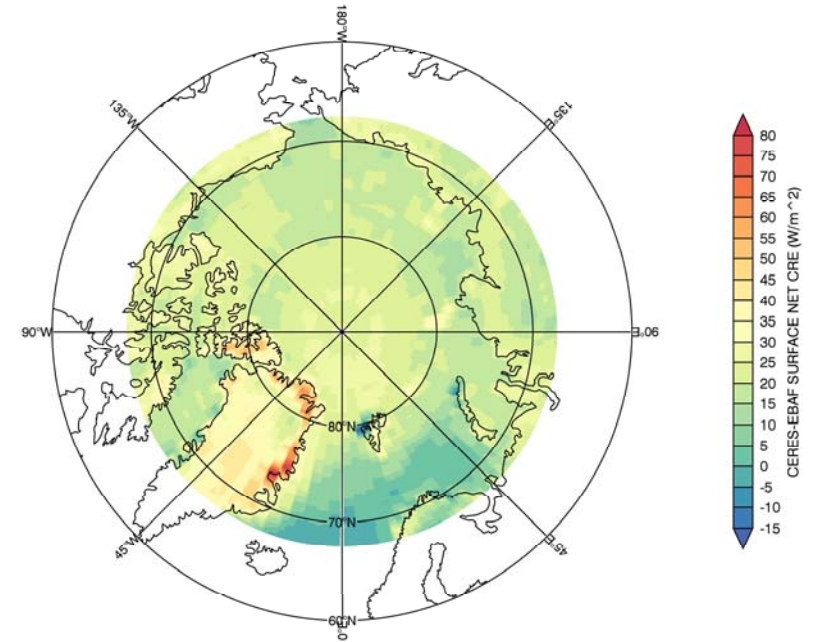
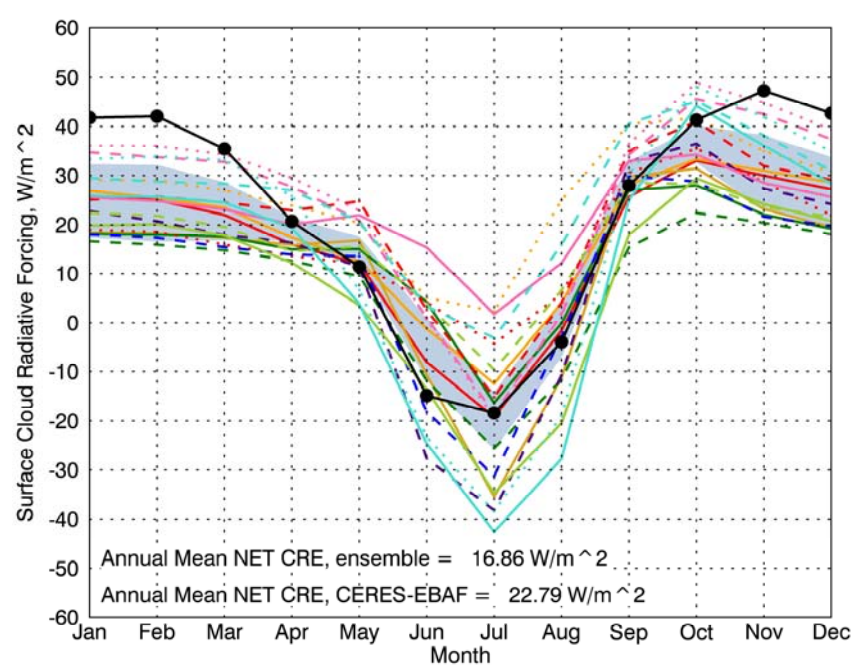
Shortwave Cloud Radiative Effect



Generally, models with higher surface albedo have a weaker SW CRE and vice versa

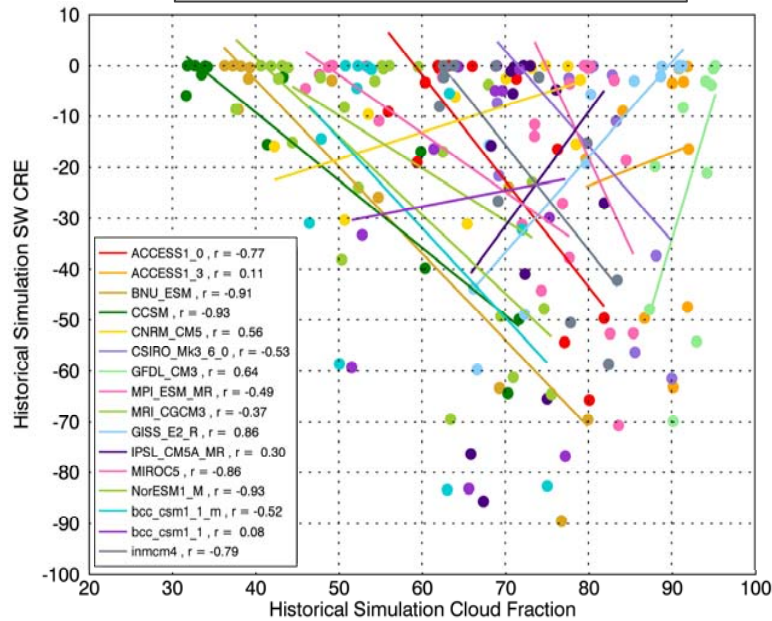


Net Cloud Radiative Effect

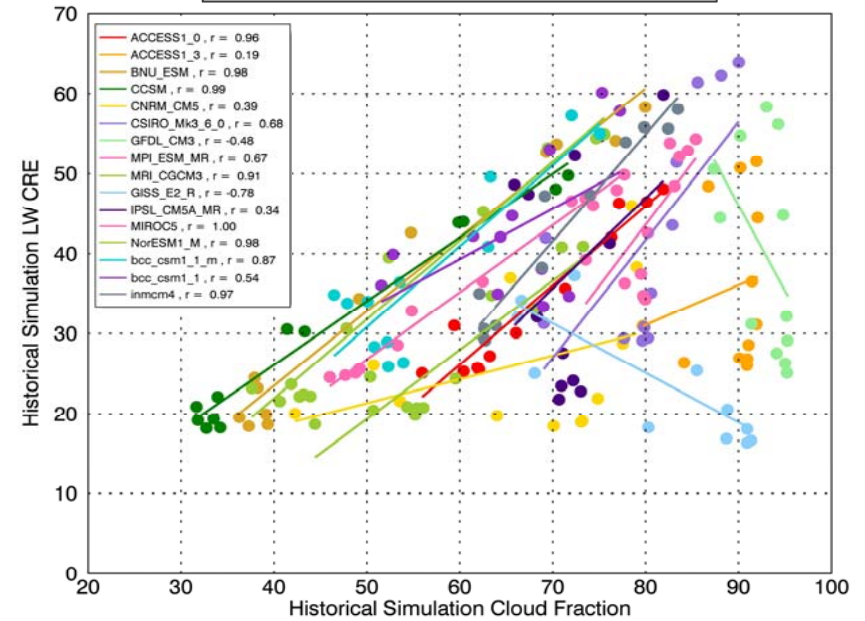


Regressions between cloud fraction and net CRE show whether a model is more strongly forced by a **cloud albedo effect** or a **cloud greenhouse effect**

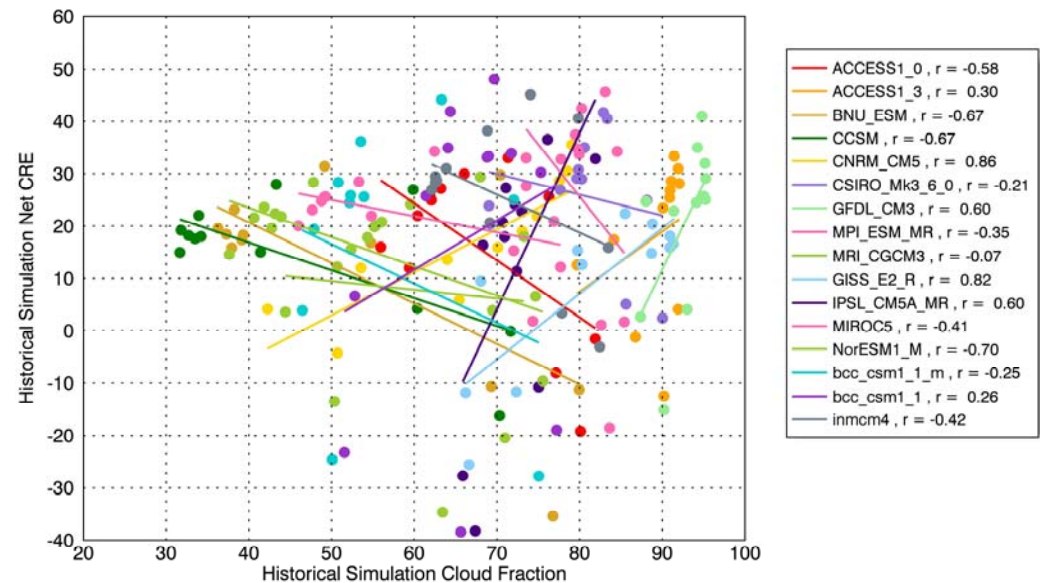
SW CRE vs Cloud Fraction



LW CRE vs Cloud Fraction



Net CRE is the result
of adding the
longwave and
shortwave forcings

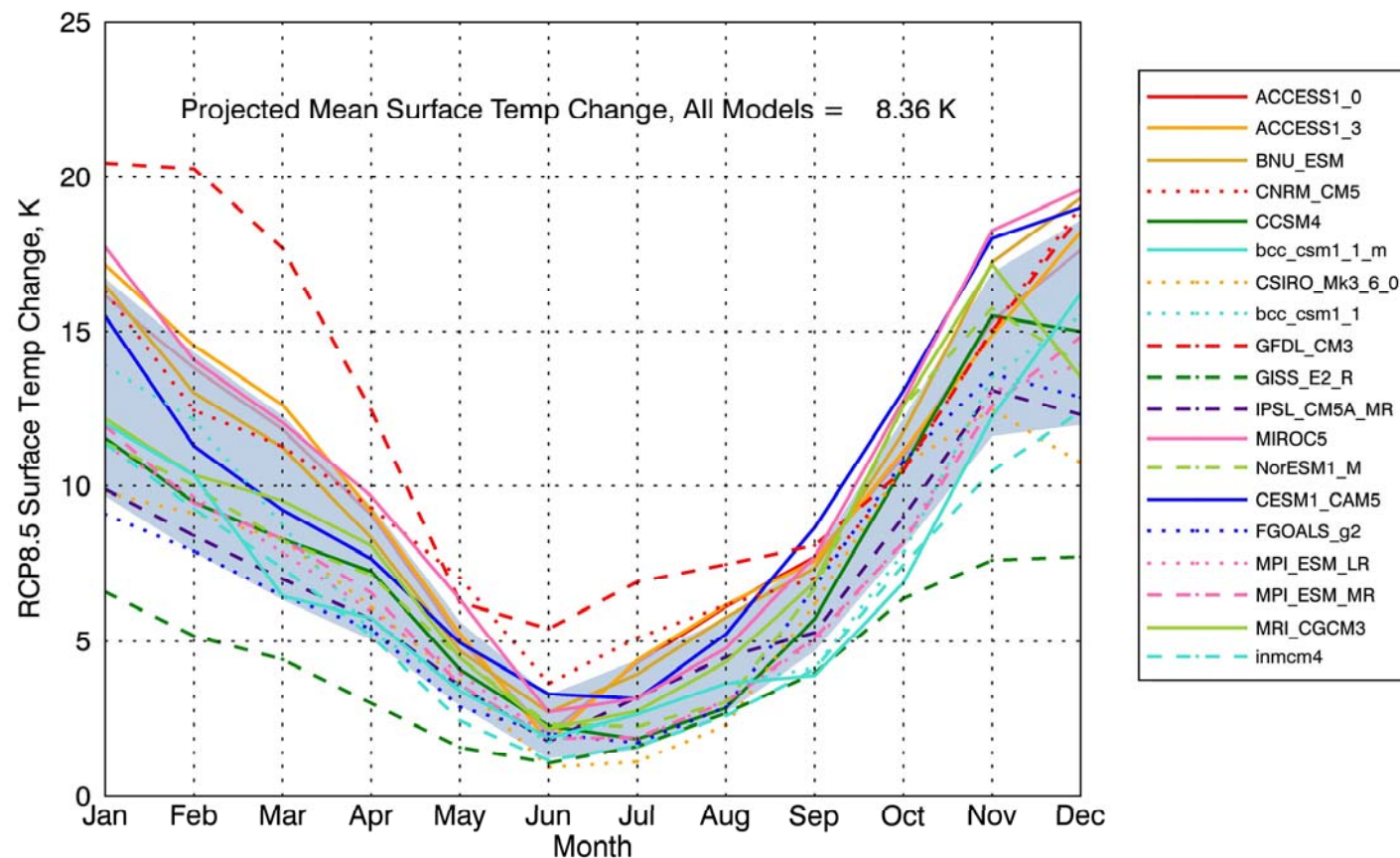


How will Arctic surface temperature change in the future?

Future surface temperature is obtained using the RCP 8.5 simulation (**R**adiative **C**oncentration **P**athway 8.5, a projection dataset with an 8.5 W/m² forcing)

RCP 8.5 runs from 2006 to 2100. Temperature change is calculated as follows:

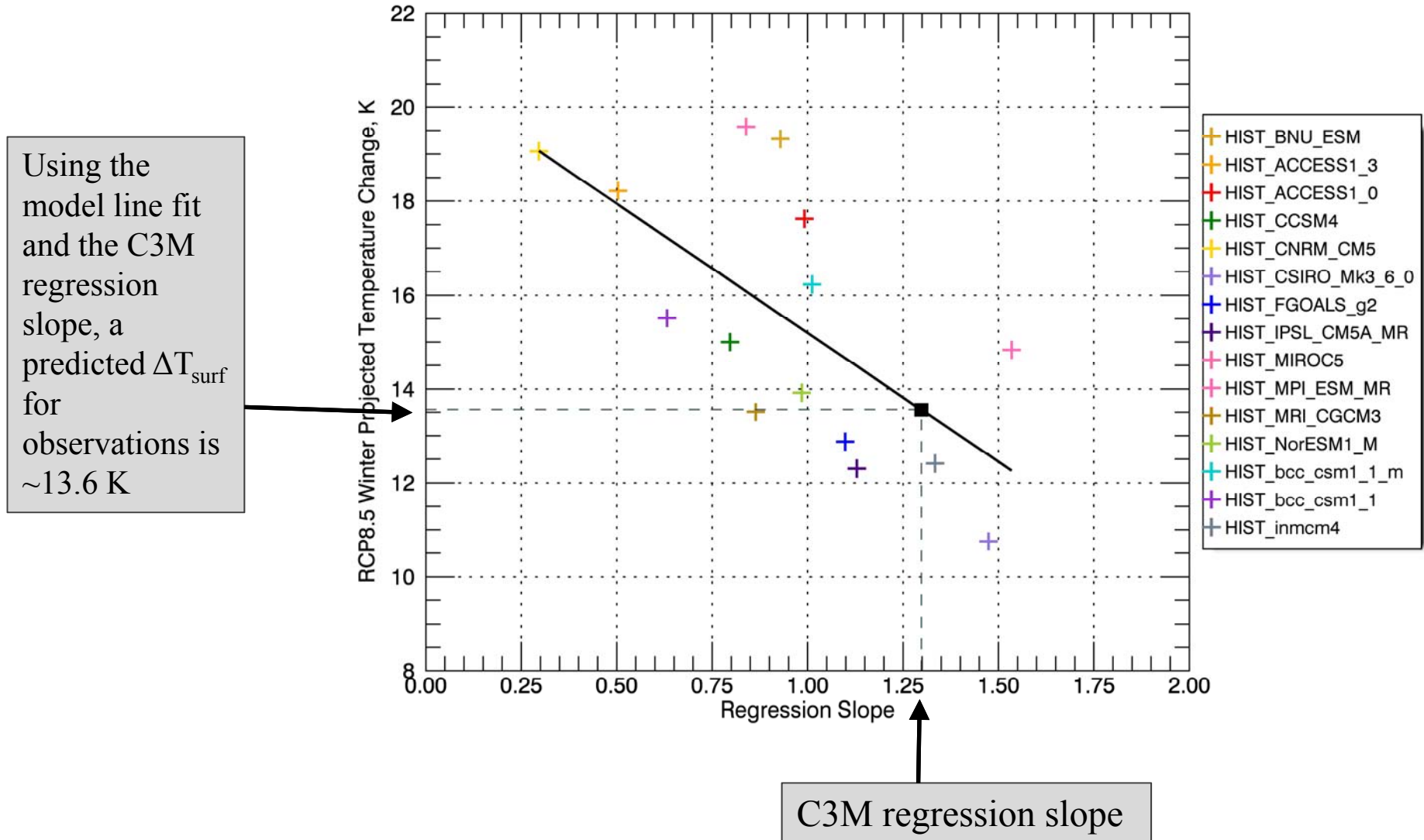
$$\Delta T_{\text{surf}} = \text{Mean } T_{\text{surf}} \text{ for the last 20 years of the simulation} - \text{Mean } T_{\text{surf}} \text{ for the first 20 years of the simulation}$$



(grey shaded region is the ensemble mean +/- one standard deviation)

The sensitivity of a model to changes in clouds is correlated to projected surface temperature change

The slope of the regression line from the δN vs δLW CRE is compared to projected ΔT_{surf} for CMIP5 models and C3M observations



Questions?

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